

Sensitivity on earth core and mantle densities using atmospheric neutrinos

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The Earth mass

Eratostene
(273-192 b.C.)



First measure of the Earth circumference
(and then of the Earth radius)

Galilei
(1564-1642)



Determination of the acceleration due to
gravity

Newton
(1642-1727)



Formulation of the law of gravitation

Cavendish
(1731-1810)



Determination of G_N



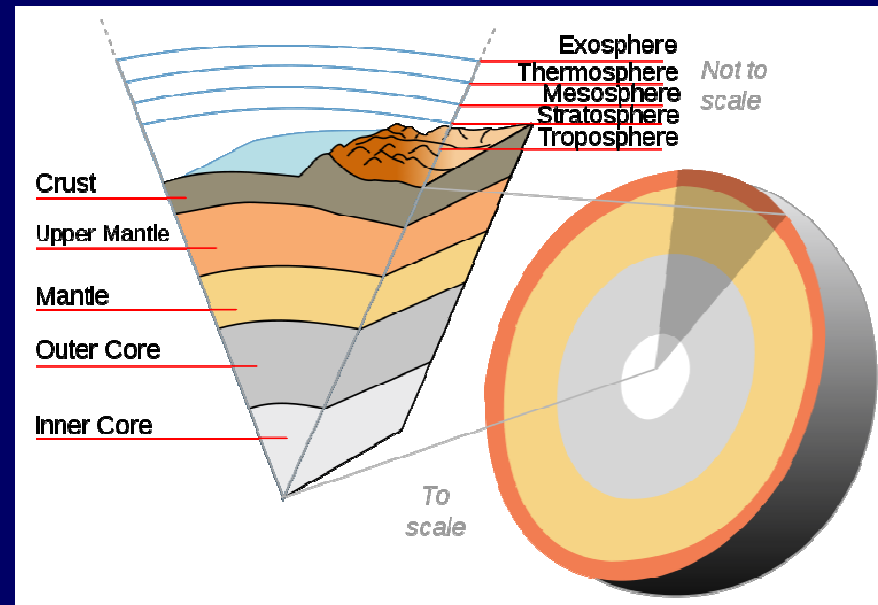
Calculation of the Earth mass and then
of its average density:

$$\rho_{av} \sim 5.5 \text{ g cm}^{-3}$$

The Earth internal structure

On the other side, the Earth crust density is about $2.7\text{-}2.8\text{ g cm}^{-3}$ (direct observations arrive to $\sim 20\text{ km}$). Information from samples brought to the surface by volcanic activity and by measuring the travel times of earthquake waves to seismograph stations. It is found that:

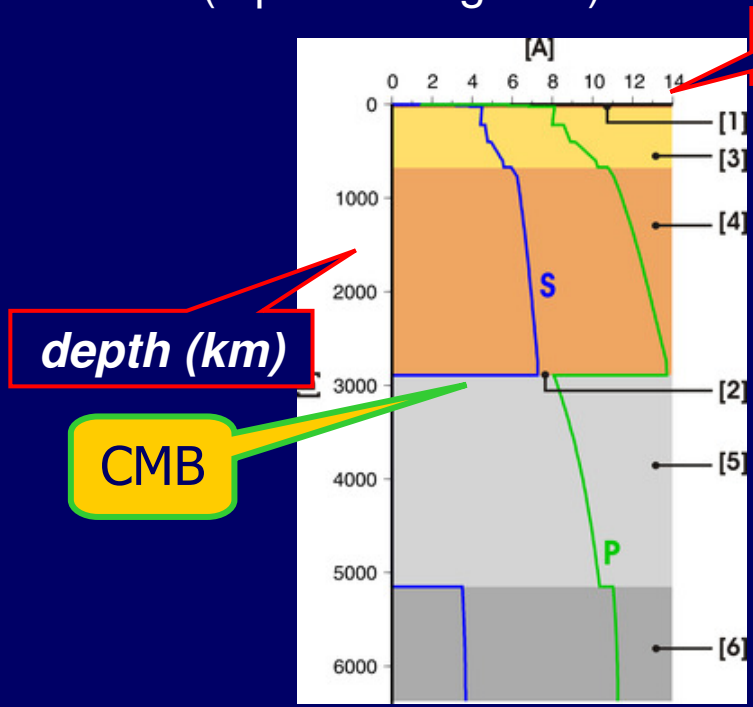
- the velocity generally increases gradually with depth in Earth, due to increasing pressure and rigidity of the rocks
- however, there are abrupt velocity changes at certain depths, indicating layering



The utility of earthquakes...

In solids, the P waves generally travel almost twice as fast as S waves and can travel through any type of material. S waves can travel only through solids, as fluids (liquids and gases) do not support shear stresses.

$$V_p = \sqrt{\frac{k + \frac{4}{3}\mu}{\rho}} \quad V_s = \sqrt{\frac{\mu}{\rho}}$$



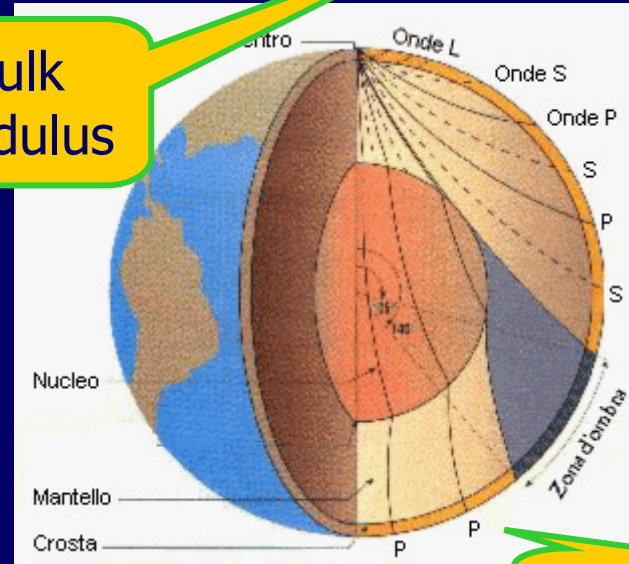
$v \text{ (km s}^{-1}\text{)}$

depth (km)

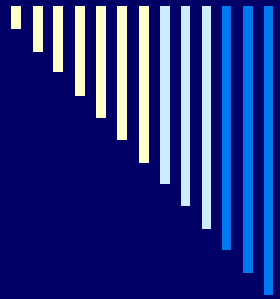
CMB

bulk modulus

modulus of rigidity

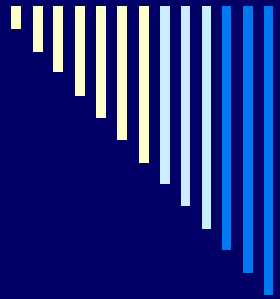


PKP (P')



Limits of seismic tomography

- Global seismic tomography is limited by the irregularity in time and space of the source, and by the incomplete coverage of recording stations. The primary source is earthquakes, which are impossible to predict and only occur at certain locations around the world. In addition, the global coverage of recording stations is limited due to economic and political reasons. Because of these limitations, seismologists must work with data that contains crucial gaps.
- Free-oscillation data only reveal 1-dimensional structure.



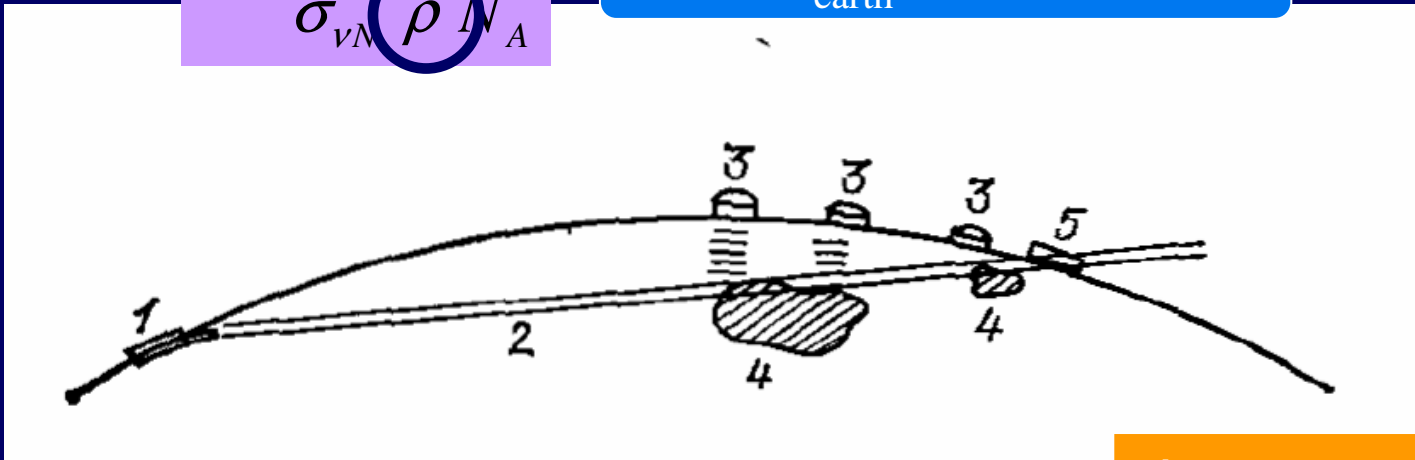
Limits of seismic tomography

- Global seismic tomography is limited by the irregularity in time and space of earthquakes, the incomplete coverage of seismic stations, and the fact that earthquakes only occur at certain depths and locations. The global coverage of seismic stations is limited by political boundaries and must be improved.
- Although this information is more precise than what we can realistically expect from neutrino radiography in the near future, aspects of the global structure of the earth require confirmation.
- Free-oscillation data only reveal 1-dimensional structure.

Earth radiography by neutrinos

$$\lambda = \frac{1}{\sigma_{\nu N} \rho N_A}$$

$$\lambda \sim 2 R_{\text{earth}} \Leftrightarrow E \sim 20 \text{ TeV}$$



Askaryan, Usp. Fiz. Nauk 144, 523 (1984)

Phys. Rep. 99, 341 (1983)

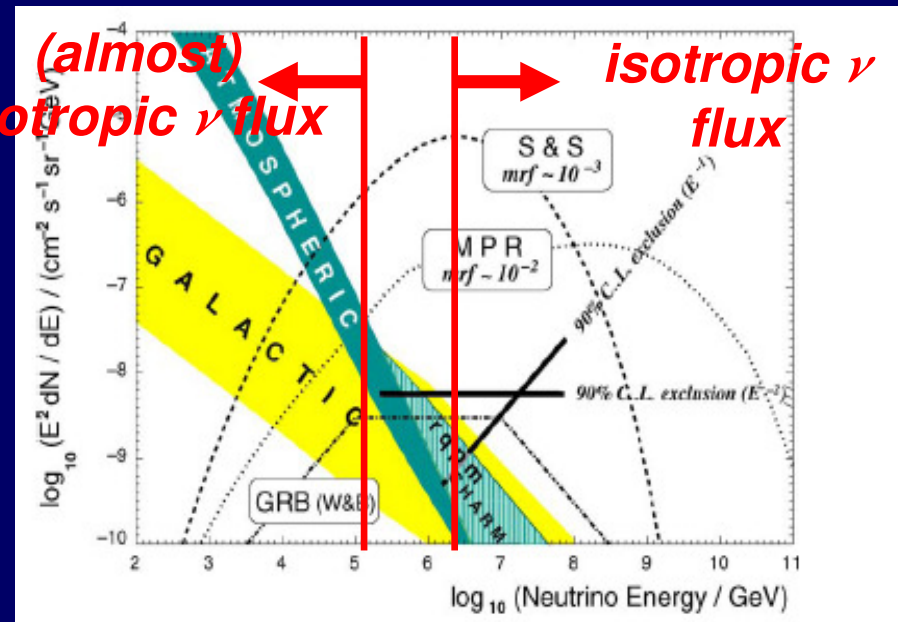
First proposal to use neutrino beams for Earth radiography by De Rujula, Glashow, Wilson, and Charpak in 1983: a neutrino moving in rock produces a shower which ionizes the medium and generates an acoustic signal. Moreover, the muons accompanying the neutrino beam can be detected at the point of emergence from the Earth.

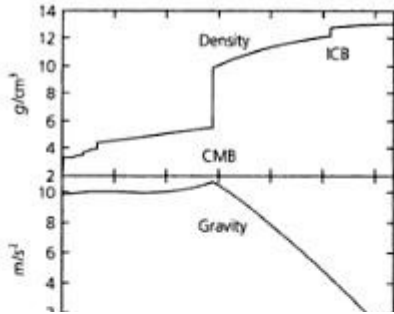
Which ν 's for Earth radiography?

Neutrinos

Neutrinos are one of the components of cosmic radiation. The atmospheric and the extragalactic contributions, which are mainly isotropic, dominates on the galactic component in two energy regimes. In particular, for $E \lesssim 10^5$ GeV the statistics of atmospheric ν is larger than that expected from other cosmic sources.

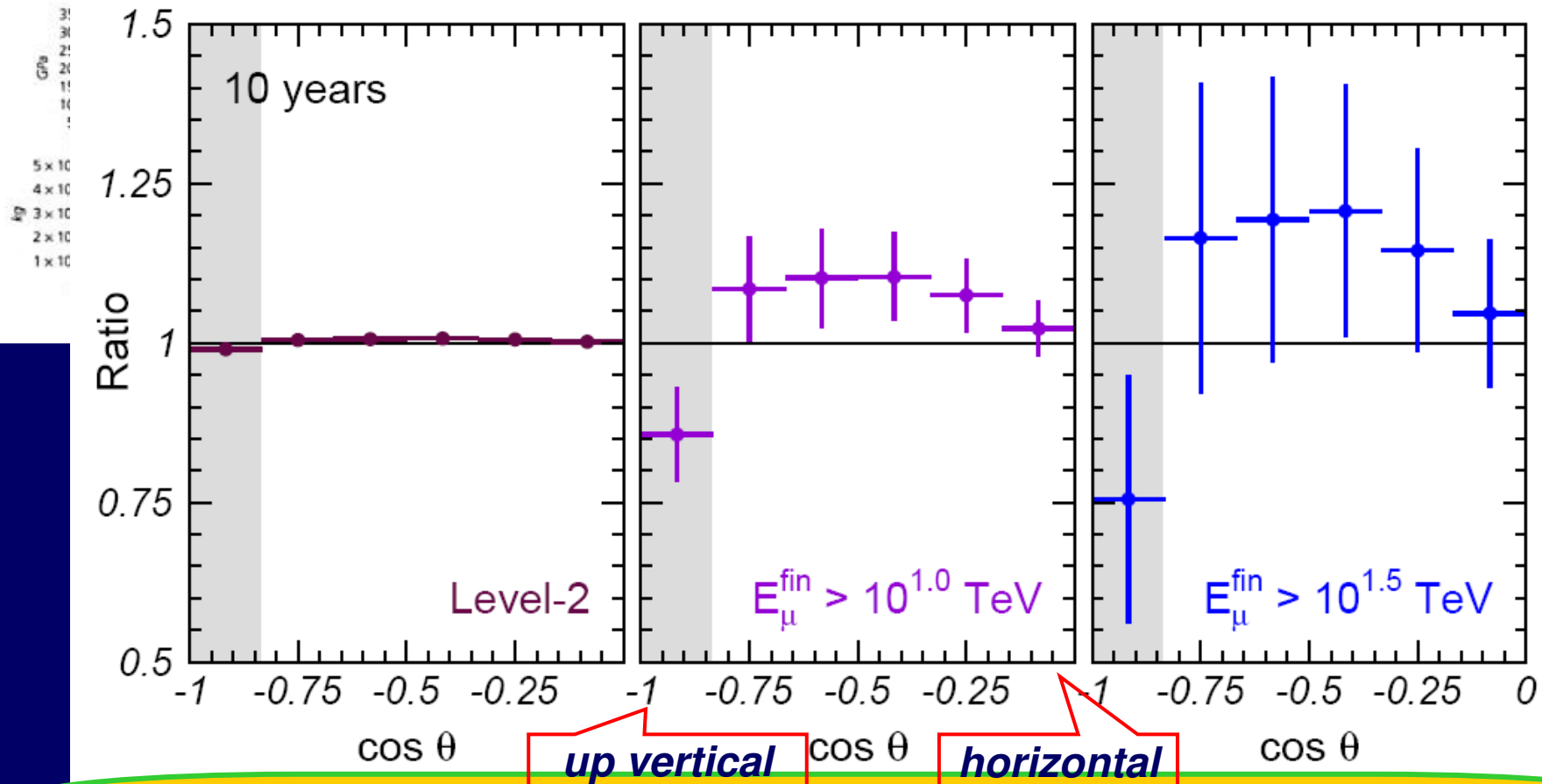
Astrop. Phys. 20 (2004) 507





PREM versus homogeneous

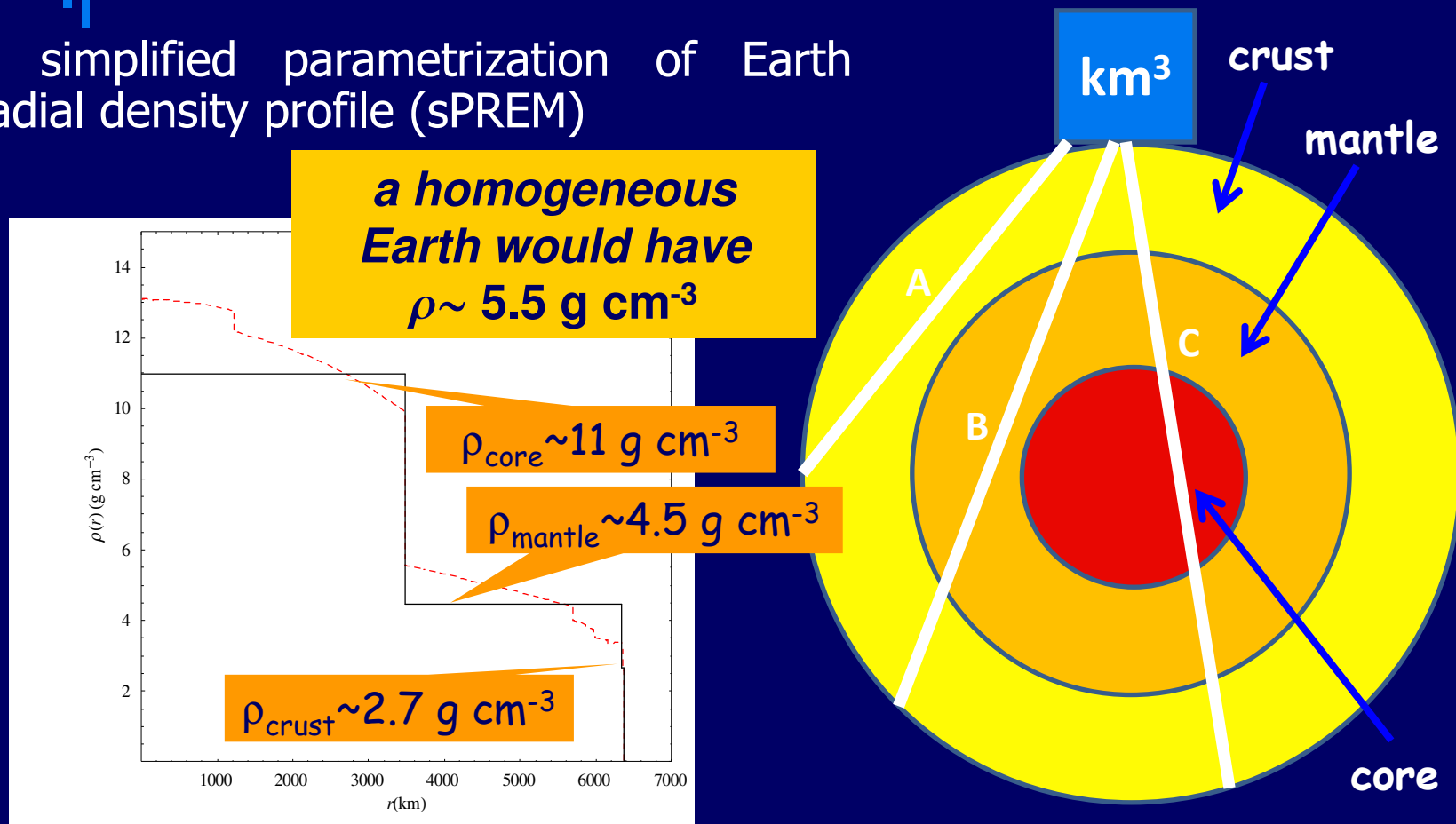
Gonzales-Garcia, Halzen, Maltoni, Tanaka, PRL100 (2008) 061802



For high E_{th} the attenuation factor due to the earth density becomes relevant and the number of detected events gives information on the earth density.

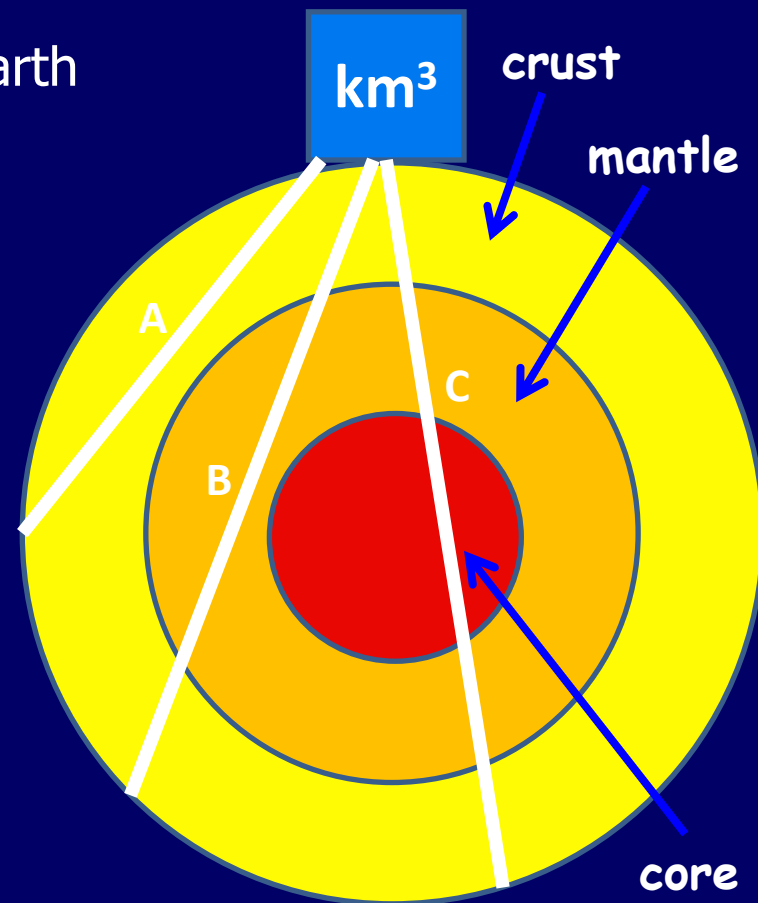
Neutrinos through the Earth

- a simplified parametrization of Earth radial density profile (sPREM)



Neutrinos through the Earth

- a simplified parametrization of Earth radial density profile (sPREM)
- three different type of tracks crossing the fiducial volume (described by a Digital Elevation Map)
- neutrinos injected at one side of the tracks according to known spectrum

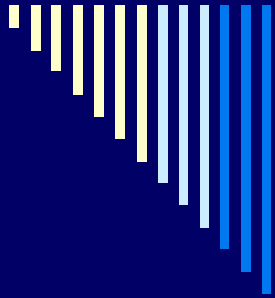


Honda et al., PR D75:043006 (2007)



Summary of simulation details

- e and μ (anti)neutrinos injected according to the atmospheric ν flux in the range $1\div 10^2$ TeV (Honda et al., 2007). Negligible oscillation \rightarrow no τ contribution
- CC neutrino interaction and neutrino regeneration by NC processes
- μ energy loss in matter (ionization, bremsstrahlung, pair production, nuclear interaction)
- detectable events: *track* and *contained* events
- energy thresholds of 1 TeV
- no details of the experimental apparatus, except for the request of a minimal track length of 300 m in the NT



Likelihood analysis

$$4 \leq \rho_m / (\text{g cm}^{-3}) \leq 5$$

$$9 \leq \rho_c / (\text{g cm}^{-3}) \leq 12$$

We consider 5 angular bins for the interval $\cos\theta = 1$ (upgoing) to $\cos\theta = 0$ (horizontal) and make the analysis integrating the muons at different energies in the case $E_{\text{th}} = 1$ TeV.

Observables N_i produced for a grid of 20 theoretical models of densities. Then, likelihood analysis with likelihood function $L \propto e^{-\chi^2/2}$ and

overall uncertainty on ϕ and σ

$$\Delta\xi = 0.25$$

$$\chi^2(\rho_m, \rho_c, \xi, \eta) = \sum_{i=1}^5 \frac{[N_i(\rho_m, \rho_c)(1 + \xi)(1 - \eta \langle \cos \vartheta \rangle_i) - N_i^0]^2}{N_i^0} + \left(\frac{\xi}{\Delta\xi}\right)^2 + \left(\frac{\eta}{\Delta\eta}\right)^2$$

uncertainty between h. and v. ϕ

$$\Delta\eta = 0.05$$

Likelihood analysis

$E_{th} = 1 \text{ TeV}$

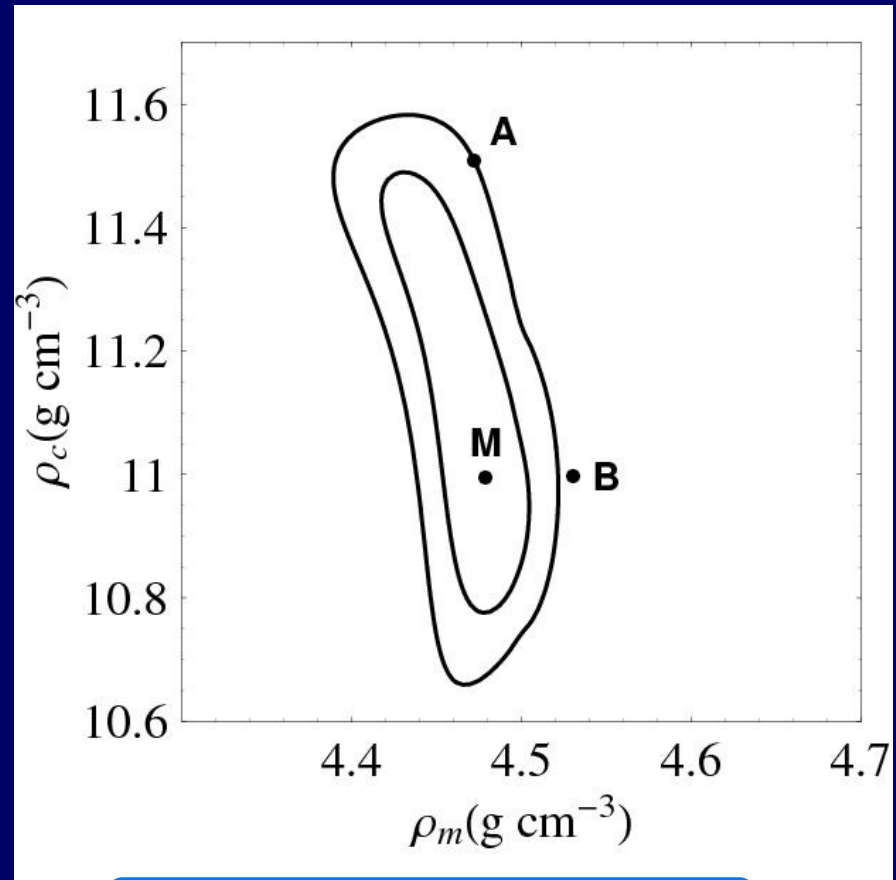
**2% (5%) uncertainty
on ρ_m (ρ_c) at 2σ**

$$\rho_m = 4.47^{+0.02}_{-0.03} \left(\begin{matrix} +0.04 \\ -0.06 \end{matrix} \right) \text{ g cm}^{-3}$$

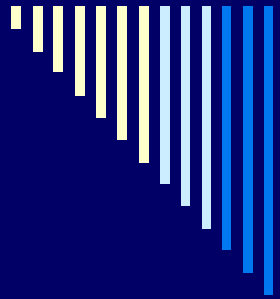
$$\rho_c = 11.0^{+0.3}_{-0.1} \left(\begin{matrix} +0.5 \\ -0.2 \end{matrix} \right) \text{ g cm}^{-3}$$

$$R_c = 3440 \pm 30 \left(\begin{matrix} +70 \\ -50 \end{matrix} \right) \text{ km}$$

$\cos \vartheta$	sPREM (M)	A	B
[0, 0.2]	113436	113860	112876
[0.2, 0.4]	72393	75456	73981
[0.4, 0.6]	47334	48142	47790
[0.6, 0.8]	34105	34144	33503
[0.8, 1.0]	26781	27392	26780



10 years of observations



Conclusions

- study of the sensitivity of a NT to Earth interior for a simplified Earth model (sPREM)
- 2% (5%) uncertainties (at 2σ level) on ρ_m (ρ_c) for 10 years of observations and $E_{th}=1$ TeV
- no details of the experimental apparatus
- low number of model parameters \Rightarrow good level of sensitivity in their determination